

# Groundwater Occurrence and the Dissolution of Salt at the WIPP Radioactive Waste Repository Site

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## Introduction

The Waste Isolation Pilot Plant (WIPP), located about 25 miles east of Carlsbad, in southeastern New Mexico, is slated to be the first deep geologic repository for permanent disposal of radioactive wastes in the United States. The repository will be located in bedded salt of Permian (225 m.y. B.P.) age, at a depth of 655 m below the ground surface. The present mission of WIPP calls for a permanent disposal of approximately 170,000 cubic meters of defense transuranic (TRU) wastes and for temporary, retrievable emplacement of 4.25 cubic meters of experimental high level wastes. The site will not be licensed by the Nuclear Regulatory Commission (NRC) but will comply with the Environmental Protection Agency (EPA) standards and all other federal and state regulations.

The site evaluation work for the WIPP has mostly been conducted since 1975 by Sandia Laboratories and the U.S. Geological Survey for the U.S. Department of Energy. On June 1, 1989, the Department of Energy (DOE) declared the site to be acceptable, and announced its decision to proceed with the construction of the repository. Drilling of two shafts to a depth of 655 m and excavation of more than 9000 m of drifts has already been completed as part of the site and preliminary design validation (SPDV) program. Underground excavation is currently in progress to excavate the "rooms" where rock mechanics and heater experiments without radioactive materials will be conducted until 1988. The radioactive waste is scheduled to start arriving in 1989. The decision on whether or not to retrieve the TRU waste will be made 5 years after the first emplacement at WIPP. According to present schedule, the site will be decommissioned in 2006. This date primarily reflects the anticipated life of the mine shafts at WIPP.

The New Mexico Environmental Evaluation Group (EEG) conducts an independent review of all maps for the WIPP project for the State of New Mexico. After a careful review of the results of site investigation, EEG [Noll *et al.*, 1983] has concluded that the site selected for WIPP has been characterized in sufficient detail to warrant confidence in its suitability, although additional work remains to be done to answer fully some remaining questions. These questions concern the hydrology and the mechanics and rate of salt dissolution. The DOE is committed to resolve these questions satisfactorily before bringing any radioactive waste to WIPP. The discussion that follows provides the background and outlines the need for further understanding of these questions.

### Geologic Setting

The WIPP site is situated in the northern part of the Beekmantown Formation, which is

part of the Delaware Basin, which is a sub-basin of the well-known Permian Basin of the southwestern United States. The Delaware Basin is bounded by a Permian reef, known as the Capitan Reef (Figure 1). The basin contains about 4300 m of Paleozoic sedimentary rocks overlying the Pre-Cambrian basement. The upper 1250 m consist of a sedimentary sequence belonging to the Ochoan Series (Upper Permian), the lower 1000 m of which consist of the three evaporite forma-

formations, from oldest to youngest, are Castle, Little Salado, and Rustler, respectively (Figure 2). Underlying the Ochoan series formations is the Delaware Mountain Group (DMG) formations which form the floor of the Delaware Basin evaporite sequence. The total thickness of the DMG is about 1800 m, but its upper

The site lies on a generally flat plain covered with sand and caliche and desert bushes. Two actively developing solution-erosion features lie west and east of the site. The one to the west, Nash Draw, is a dog-bone shaped subsidence feature 10–20 km wide in the

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east-west direction and about 30 km long (north-south). The eastern edge of Naah Draw lies 5 km west of the edge of the WIPP site (Figure 3). (Figure 3 shows the original boundary of WIPP, the outer fringe of which was abandoned by DOE in 1985. The presently planned repository is a 100 acre area south of ERDA-9. This area may be extended later but will remain within the inner octagonal area known as Zone 11. Outside this is Zone 111, where no surface mining or non-WIPP drilling will be allowed. A fenced area at the center (not shown in Figure 3) will house the administrative buildings and the shafts and is designated Zone 1.)

## Geohydrology

The Ruisder and the Bell Canyon formations contain poor quality groundwater under confined conditions. In addition, pressurized brine has been encountered in the upper part of the Castle Formation in several wells near the WIPP site (Figure 3). Since groundwater is the most likely mechanism for bringing radioactive waste from the WIPP repository to the biosphere, a detailed knowledge of the regional site specific groundwater conditions is necessary to assess the radiological consequences of breach of the repository. A summary of the characteristics of various aquifers at the site follows.

### Bell Canyon Formation

Over the Delaware Basin as a whole, groundwater in the Bell Canyon Formation flows from west to east. In the vicinity of the WIPP site, flow is to the northeast. The unit is recharged along the western edge of the Delaware Basin, 90 km west of the site, where the Bell Canyon crops out. The unit appears to discharge into the Capitan Reef to the east (Figure 2).

The salinity of the Bell Canyon water increases from west to east. In the vicinity of the site, the total dissolved solids in the Bell Canyon range from 180 to 270 g/l. The water in the Bell Canyon is under artesian pressure and the potentiometric surface corrected to fresh water potential rises to ground level at the WIPP site.

Hydraulic head data are sparse and of questionable accuracy. In the vicinity of the WIPP site, only three boreholes have so far been used to collect Bell Canyon head data. A direct measurement of hydraulic head is available at only one well. At the other two, down hole pressure and fluid density were used to calculate the hydraulic head. It is planned to test this formation in three more boreholes during 1984-1985.

### Geologic Setting

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The site lies on a generally flat plain covered with sand and caliche and desert bushes. Two actively developing solution-erosion features lie west and east of the site. The one to the west, Nash Draw, is a dog-bone shaped subsidence feature 10–20 km wide in the

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61, is that the two boreholes intercepted separate reservoirs. However, although it would be incorrect to conclude that each of the 13 occurrences represents a separate brine reservoir, there is no indication that a brine "aquifer" exists in the Castile.

The brines are of concern because they are under anomalously high pressures. In WIPP-12, the pressure of the brine reservoir encountered 930 in below the ground surface) is sufficient to support a column of brine extending more than 300 m above the ground surface. If a brine reservoir exists beneath the repository, then the potential for bringing radioactive material to the surface exists. EEO [Chernoff, 1982; Bard, 1982] has calculated the maximum radiation dose to an individual from a brine reservoir scenario to be below the limits that EPA Protective Action Guides recommend for low probability accidents. These calculations may change substantially if the type of radioactive wastes to be buried at WIPP is changed from that presently planned.

### Rustler Formation

The Rusler Formation directly overlies the Salado Formation and contains three recognized fluid bearing zones. From stratigraphically lowest to highest, these are the Rusler-Salado contact residuum, the Culebra dolomite, and the Magenta idolomite. The transmissivity of the Culebra is highest, followed by the Magenta and the Rusler-Salado contact. The water quality is highly variable within each unit. The total dissolved solids concentration is lowest in the Magenta (the uppermost unit) and highest in the Rusler-Salado (the lowestmost unit). Nearly all the water in the Rusler has TDS concentrations greater than 10,000 mg/L. The exception is the Culebra water south-southwest of the center of the WIPP site where the TDS is about 3,300 mg/L.

All three units probably discharge into the Pecos River 25 km southwest, near a bend in the river known as the Malaga Bend (Figure 1). The recharge areas are identified rather imprecisely as being upgradient of the measured hydraulic heads about 15–25 km north of the WIPP site. At the WIPP site the three units are separated but are probably interconnected in Nash Draw west and southwest of the site.

Of the three Rusler onits, the Magenta and Gulebra are of prime concern because they extend over the WIPP site, whereas the Rusler-Salado contact zone mainly produces water west of the WIPP site [*Coubatz*, 1983; *Mirrer*, 1983]. The majority of the testing in the Rusler has concentrated on the Gulebra because it is more transmissive than the Magenta and therefore better suited for analyzing the worst case bounding breach scenario.

The groundwater in the Magenta flow west from the site toward Nash Draw and then to the Pecos River near Malaga Bend. It and near Nash Draw the Magenta and Culebra appear to be hydraulically connected because the hydraulic head difference, over 30 m at the WIPP site, is only a few meters in Nash Draw. The direction of groundwater flow in the Culebra indicated by the potentiometric surface [Mercer, 1983] is south or southeast; for 6–8 km and then to the west toward Malaga Bend. This apparent flow path is inconsistent with the observed chemistry, marked change in the water chemistry in the Culebra aquifer occurs about 6.5 km south-

west of the center of the site. The total dissolved solids to the southwest is much less than to the north. In addition, the water to the southwest is dominated by  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ , and  $\text{SO}_4^{2-}$ , whereas to the north the water chemistry is dominated by  $\text{Na}^+$  and  $\text{Cl}^-$ . Comparing the subsets of Culebra water, the total  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ , and  $\text{SO}_4^{2-}$  differ by a factor of 2 or 3, whereas the  $\text{Cl}^-$  differs by factors of 8–1000. Therefore, a simple mixing of high TDS water with fresh recharge from precipitation does not explain the observed Culebra water chemistry.

### The Capitan Reef

The Capitan Reef comprises one of the most productive freshwater aquifers in south eastern New Mexico. It surrounds the Delaware Basin in a horseshoe shape (Figure 1). At its closest point, it lies 16 km to the north of the WIPP site. Water that flowed eastward from this aquifer into the Bell Canyon combined with surface recharge into the Ochoan evaporite sequence of the Delaware Basin must have been responsible for the dissolution of salt in the basin. Whether or not this is an active process at the present time is a matter of controversy and a great deal of speculation. Some [e.g., *Andersson and Kirkland, 1980; Andersson, 1981*] believe that the Capitan Reef acts as an recharge source to the west and a discharge sink to the east and has thus facilitated the dissolution of salt in the evaporites at depth in the Delaware Basin. This point is further discussed below.

### Extent of Salt Dissolution in the Delaware Basin

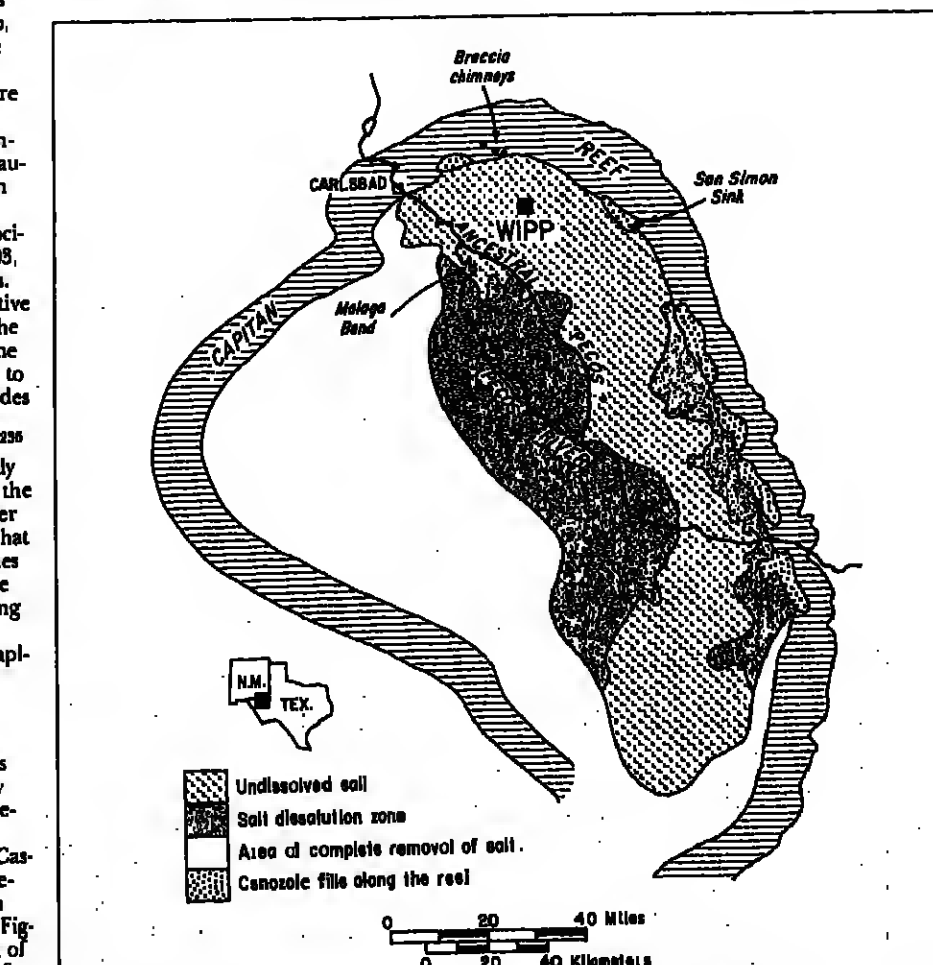
Bedded salt is a preferred medium for permanent emplacement of radioactive waste because of the favorable physical, thermal, and mechanical properties of halite. However, halite is also readily soluble in water, and most bedded salt deposits have undergone different degrees of dissolution since their deposition.

In the area of the WIPP, there are several indications that a large amount of salt from both the Rustler and Salado formations has been dissolved away. To evaluate the impact of such dissolution on the nuclear waste repository of the WIPP, it is necessary to understand the horizontal and vertical extent as well as the time and rate of salt dissolution in the Delaware Basin.

The regional dissolution of salt beds in the basin has progressed from west to east. It must have been initiated with the tilting of the basin to the east sometime during the Cenozoic era, which facilitated the injection of water from the Capitan Reef aquifer down-dip to the east into the Bell Canyon as well

dip to the east into the Seal Canyon as well as surface flow into the evaporite deposits. In the western part of the Delaware Basin, the exposed Castile, Salado, and Rustler formations no longer contain any salt. The Salado which before dissolution consists mainly of halite, is represented by a chaotic jumble of insoluble residue, less than one third the thickness. Even where the Salado is buried underground, most of the salt in it is probably represented by fractured anhydrite and breccia in the region designated as the "Area of Complete Removal of Salt" in Figure 1. The complete area of "Salt Dissolution Zone" of Figure

Article (cont. on p. 45)



**Fig. 1. Extent of removal of salt from the Salado formation in the Delaware Basin**



## Article (cont. from p. 437)

1. the borehole logs indicate dissolution in the lower Salado and upper Castile (Anderson, 1982). Barkman (1983) interprets the well logs to indicate dissolution breccia in lower Salado but not in Castile. The different interpretation stems from the difficulty in recognizing the Castile/Salado boundary.

The removal of halite by dissolution is observed in the Rustler Formation as well. Where unaffected by dissolution, the Rustler Formation is 150 m thick and contains about six discrete halite units with an aggregate thickness of about 40 m. Much of this halite has been removed over most of the WIPP site by dissolution. Directly above the selected repository area (Zones 1 and 11) of the WIPP site, all the salt in the Rustler Formation, except a layer of argillaceous halite below the Culebra aquifer, has been removed by dissolution.

## Mechanisms of Dissolution

The earliest proposal for a specific mechanism of dissolution of salt in the Delaware Basin was made by Lee (1925). His mechanism of "solution and fill" postulates infiltration of rain water collected in arroyos into the fractures of soluble rock. This results in the development of a drainage system a few feet below the surface, into which surface debris is carried by subsequent storms. As the gradient of the drainage system increases, headward cutting results. According to Bachman (1980), this process is currently active in Nash Draw, west of the WIPP site.

As stated above, salt has been removed from the Rustler Formation, which lies between 170 and 260 m below the ground surface at the WIPP site. The lowest affected zone is progressively deeper to the west. In the western part of the WIPP site and in Nash Draw, the top of Salado has also been affected by dissolution, and the permeable residuum thus formed contains brine at the Rustler/Salado interface. There are salt lakes in the southern part of Nash Draw and there are saline seeps along the Pecos River near Malaga Bend, about 25 km southwest of the WIPP site. These are thought to be the discharge points for the brine produced from the dissolution of Rustler halite.

A satisfactory explanation of the dissolution and removal of salt from the deeper strata is more problematical. There are at least three different schools of thought concerning the absence of halite in the lower Salado and Castile formations. On the basis of interpretation of acoustic logs from a large number of wells in the Delaware Basin, Anderson (1982) has concluded that about 50% of the salt in the Salado and Castile formations has been removed by dissolution, with as much as 70% of the original salt removed from the lower Salado horizon in the basin. For the mechanism of salt removal through "deep" dissolution, Anderson (1981) invoked a "brine density flow" model, which had been proposed earlier (Anderson and Kirkland, 1980) for the formation of breccia chimneys. This mechanism requires a connection between the lower Salado and the underlying Bell Canyon aquifer through fractures in the intervening Castile Formation. It was hypothesized that surface recharge moves into the evaporites, dissolves salt in the Salado and Castile formations, and the resulting brine sinks into the underlying aquifer. Thus, the postulated mechanism would continue as long as the supply of undersaturated water basins and the fractured pathway remains unclogged. Wood et al. (1982) studied the potential dissolution mechanisms of diffusion and convection from the halite zones of Castile and Salado to the Bell Canyon and the Capitan Reef aquifers based upon a range of reported values for the hydrologic and geochemical parameters which influence salt removal. They concluded that the removal of dissolved salt through the Bell Canyon can take place only at a very slow rate, which would be grossly insufficient for removing approximately  $7 \times 10^{12}$  m<sup>3</sup> of salt from lower Salado in 1.5 million years, or about 4.7 million m<sup>3</sup> of salt per year (as estimated by Anderson, 1981). Even using the most conservative reported values for the hydraulic conduc-

tivity of the Bell Canyon, Neill et al. (1983) estimated that at most 20–50 times less salt could have been transported through the Bell Canyon than is estimated to have been removed.

The second explanation for the missing salt is provided by Barkman (1983). According to him, major dissolution of the evaporites in the Delaware Basin has been restricted to areas where the Pecos River and its tributaries have initiated karst systems, or to limited areas which overlie the Capitan Reef aquifer. Figure 1 shows the path of an ancestral Pecos River, east of the present day Pecos, as postulated by Barkman (1983) on the basis of river gravel deposits left by this river. Barkman (1983) believes that this ancient system was responsible for the development of an extensive karst terrain now seen east of the present day river. The salt beds of the lower Salado Formation were selectively dissolved during the Cenozoic time as a result of a dissolution front which was perched on the upper anhydrite in the Castile Formation. Barkman (1983) further states that the Tertiary and Pleistocene hydrologic conditions no longer exist except along the present Pecos River channel, and therefore the probability of further dissolution in the proximity of the WIPP repository horizon is remote.

The third explanation for the missing salt in the lower Salado is that much of the missing salt simply represents a facies change or removal during a much earlier time soon after deposition. By drawing a composite isopach map of the Castile and lower Salado Formation, Lambert (1983) has shown that the "missing halite" areas commonly result in little or no departure from regional thickness of Castile and lower Salado evaporites. He therefore ascribes the observation of missing salt to several factors other than Cenozoic dissolution (e.g., depositional heterogeneities, perturbation of original bed thicknesses by localized deformation, and ambiguous identification of members or marker beds).

A serious problem encountered in explaining the removal of salt through dissolution is the disposition of the resulting brine. Anderson's brine density model requires the removal of brine through the Bell Canyon aquifer; but, as mentioned above, the Bell Canyon is thought to be incapable of acting as a sink for the large amount of brine produced from dissolution. Barkman (1983) has not even addressed the question of the disposal of solution brine. Lambert (1983) also has acknowledged the difficulty in postulating a viable sink for his strata-bound dissolution model.

## Time and Rate of Dissolution

Anderson (1981, 1982) maintains that most of the dissolution and removal of halite from the Salado and Castile evaporite beds in the Delaware Basin has occurred since the tilting of the basin to the east during the latter part of the Cenozoic, probably 4–8 m.y. ago. According to him, this event exposed the tilted limestone of the Capitan Reef as well as the sandstones of the Bell Canyon Formation which acted as suppliers of meteoric water to the lower part of the evaporites. Anderson has also used the correlation between the area of the bulk of the missing salt with the deep depressions filled with Pleistocene alluvial fill (described by Maley and Huffington (1953) as evidence for most of the dissolution having taken place during late Cenozoic.

However, there exists some clear evidence of at least some dissolution having occurred in earlier geologic times. Some dissolution through subaerial erosion occurred prior to the deposition of Salado, during the Permian time. This is evident from the erosional unconformity that truncates Halite III bed of the underlying Castile Formation. Barkman (1980) has proposed that the pre-Cenozoic times of nondeposition in the basin, mainly the Jurassic, must have been the times of extensive erosion and subsurface dissolution through circulation of meteoric water. Anderson (1981) disagrees with the concept of extensive pre-Cretaceous dissolution in the basin on the basis of the isopach maps of the middle and upper Salado showing regionally normal thickness trends extending westward to the area of present truncation. He points out that a truncation surface sufficient to al-

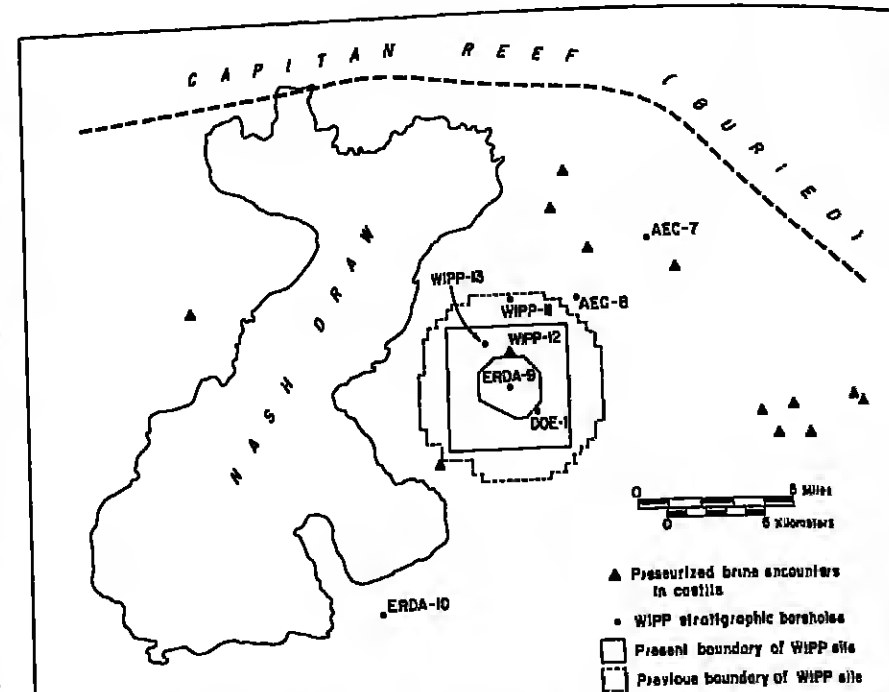


Fig. 3. Detailed map of the northern Delaware Basin showing the WIPP site, "Nash Draw" depression, boreholes encountering pressurized brine in upper Castile, and other deep stratigraphic boreholes drilled in connection with the WIPP project.

low pre-Cretaceous dissolution to reach into the Castile Formation would require a regional dip in the Cretaceous far greater than observed. Barkman (1983) believes that Tertiary and Pleistocene hydrologic conditions no longer exist except along the present Pecos River channel, and therefore the dissolution of halite in the lower Salado and Castile formations is not an active process, at least near the WIPP site.

The easiest way to determine the safety of the WIPP site from the effects of deep dissolution would be to calculate an average rate of halite removal from the Delaware Basin and the distance of the dissolution front from the site. Assuming that the edge of the Salado salt has moved from the Capitan Reef front to its present location (Figure 1) during the past 7–8 million years, Barkman and Johnson (1973) concluded that the horizontal rate of movement of the blanket dissolution front is about 10–13 km per million years. Based upon the dating of a volcanic ash layer associated with the Pleistocene Gatuna Formation, which is exposed at the ridge on the eastern margin of Nash Draw, Barkman (1980) concluded that about 60 m of subsidence has occurred in this depression during the past 800,000 years. Using this rate, Barkman (1980) calculated an average rate of 100 m per million years for the vertical dissolution.

It is more difficult to determine the rates of advance of dissolution fronts for the deep-seated variety of dissolution. Such deep-seated undermining does not necessarily leave a discrete geomorphic feature such as a scarp. Anderson's assumption that the missing salt was removed by dissolution during late Cenozoic has been questioned. According to Lambert (1983), the tilting of the basin has not been precisely dated, but only estimated by Hoya (1964) as "late Cenozoic" in the total absence of geologic time markers. The occurrence of Gatuna sediments of Pleistocene age at the top of the fills in the Maley and Huffington (1953) depressions establishes the minimum age of the fill as Pleistocene. The deep parts of these fills have not been studied to determine a maximum age for the fills, and therefore it is not reasonable to constrain the entire accumulation of the fills to the Pleistocene. However, the generally unconsolidated nature of these fills shows them to be of a recent geologic time period and their involvement in the regional salt dissolution process cannot be ruled out.

## Integrity of the WIPP Site

The question of prime importance for the WIPP nuclear waste repository is whether there is a pathway for release of radionuclides to the biosphere in case of a breach of the repository. The site investigation work conducted during the past 8 years has attempted to resolve this question. Although a great deal of information has been gathered to develop worst-case release scenarios, more work needs to be done to achieve a desirable level of assurance about the integrity of the site.

The primary potential path of breach of the repository is through the Rustler aquifers into the Pecos River. A large amount of information on the hydrologic characteristics of the Rustler Formation is available. However, the following questions remain to be answered.

1. Does water exist only in three discrete zones, or does some water move through other parts of the formation as well?
2. Are there pathways of interconnected fractures from the center of the site to the Pecos River through which transport may be much faster than the average transport time computed assuming an equivalent porous medium?
3. Are there karst channels through the Rustler?
4. What is the mechanism of removal of salt from the Rustler Formation?

5. What is a reasonable explanation for the variations in the chemical composition of Rustler water at the WIPP site, south of the site, and in Nash Draw?

The results of field testing and analyses conducted so far provide partial answers to these questions. The Department of Energy is planning to conduct additional testing and analyses over the next 2–3 years to answer these questions fully.

With respect to the front of "shallow" dissolution, the most conservative approach would be to take no credit for horizontal movement of this front, in other words to assume that the dissolution front has already moved over the site at the base of Rustler. Using the rate of vertical dissolution of 100 m per million years calculated by Barkman (1980), it would require 4.5 million years to remove 450 m of salt in the Salado above the repository horizon. Admittedly, these rates are very approximate, but they are based on conservative assumptions. There appears to be a sufficient margin of safety in the future direct breach of the repository through "shallow" dissolution of the type which has formed Nash Draw.

With respect to "deep" dissolution, as postulated by Anderson, it is very difficult to fix a rate of advance of the dissolution front. There is clear indication in the "Salt Dissolution Zone" of Figure 1 that salt is not well understood. Besides, according to Anderson (1982), the dissolution at depth may not proceed as a clearly defined "front," but the advanced effects of dissolution may be noticed along "fingers" much ahead of the dissolution front. The best way to determine the integrity of the WIPP site from this kind of dissolution is to examine the evidence from boreholes drilled at the site and surrounding it. Figure 1 shows the "deep dissolution" edges for the salt units, as interpreted by Anderson (1981). It should be noted that the WIPP site is situated in the northern part of the basin, away from the dissolution fronts. The nearest point of the dissolution edge from the WIPP site is about 25 km away to the southwest (Figure 1).

There are five boreholes at the WIPP site (WIPP-9, 11, 12, 13, and DOE-1) which have penetrated the lowermost anhydrite bed (Anhydrite-I) in the Castile Formation. These holes have been cored at selected intervals, and geophysical logs for the entire depths have been obtained. In addition, three holes have been drilled, AEC-7 and 8 to the northeast and ERDA-10 to the southwest, were drilled, cored, and logged through the Castile Formation. None of these eight boreholes (Figure 3), and none of the several hundred boreholes around the WIPP site, show any evidence of extensive dissolution. This any evidence of the fact that at least the immediate area surrounding the WIPP site has not been affected by deep dissolution. Even if the rate of advance of "deep" dissolution of Barkman (1980) (i.e., 10–13 km per million years) the site appears to be safe for the next 2 million years.

A borehole located 3.25 km north of the center of the site, drilled to assess the occurrence of potash minerals in middle Salado, encountered the Salado marker beds at elevations about 25 m below their expected occurrences. This anomalous depression has been confirmed by the logs of two other boreholes drilled from the center of the site as a possible site of deep dissolution. The Department of Energy has accepted EGG's suggestion to drill a borehole, to be called DOE-2, will be drilled and cored down to the Bell Canyon Formation. The work is planned for 1984. One more anomalous feature has been pointed out near the WIPP site. About 8 km southwest of the center of the WIPP site, the acoustic log of a well shows that 60

bed in the lowest part of the Salado Formation may be missing. Since this well is outside the WIPP boundary and is one unconfined anomaly out of a large number of wells, EEG has not made any recommendations to explore this feature further.

Approximately 3000 m of underground drifts at the selected repository level 655 m below the surface have already been excavated at the WIPP site. This excavation has been conducted to validate the site characterization under the site and preliminary design validation (SPDV) program. The thickness and continuity of strata displayed in this excavation are remarkably uniform, and there is no indication of dissolution either at the repository horizon or in cores of 15 m vertical boreholes drilled from the floor and ceiling of the excavation.

## Conclusions

There are two main areas of concern with regard to the suitability of the WIPP site. The characteristics of the water-bearing zones of the Rustler Formation should be understood very thoroughly to preclude any possibility of these zones acting as pathways for migration of radionuclides to the biosphere. In addition to hydrologic testing already completed during the past 8 years, additional drilling and field testing is planned over the next 2–3 years. These forthcoming studies include the drilling and testing of several new hydrologic wells, drilling of new wells near existing hydrologic wells to convert at least eight single wells to sets of three nested wells for tracer tests and flow tests. The study of cores and determination of hydrologic properties from the cores as well as more geochronological testing of Rustler Formation waters are also planned. The data gained from these studies will be used to refine a hydrologic and contaminant transport model of the Rustler aquifers. In addition, a water-balance study for the site will be conducted and some suspicious depressions on the site would be bored to see

whether they are alluvial dolines. Another study will try to answer the question of the mechanism of salt removal from the Rustler Formation. These studies will help to resolve the question concerning the possibility of karst conditions in the Rustler Formation. The second area of concern is the effect of dissolution of salt on the integrity of the repository. The Salado Formation does not appear to have been affected in and around the WIPP site by past dissolution. The suspicion of an area of point-source dissolution from below, located 3.25 km north of the center of the site, will be investigated. Although strata-bound dissolution of the Rustler salt occurs across the WIPP site, such dissolution does not seem to have affected the top of the Salado Formation at the WIPP site. However, the collapsed depression of Nash Draw is only about 6.5 km west of the center of the site. It is therefore important to understand thoroughly the mechanisms of removal of salt from the Rustler Formation.

The mission of WIPP calls for the permanent emplacement of transuranic waste which would be reduced to a level of radioactivity of natural uranium ore in less than 100,000 years. The site selected for WIPP has been characterized sufficiently to enable the analyses of worst case scenarios of breach of the repository and consequent release of radioactivity. However, a few gaps remain in our knowledge of the geologic and hydrologic characteristics of the site which relate to the transport of radioactivity to the biosphere in the event of a breach. Additional field work to close these gaps is being performed currently and will be completed before the waste is brought to the site in 1989. If the additional work indicates a possibility of release of hazardous quantities of radioactivity to the biosphere, the EEG will recommend additional engineered barriers or investigations for a new site.

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Cover. This photograph, looking eastward to the new lobe on the west side of Mount St. Helens' composite dome, was taken on June 18, early in a recent extrusion episode. By July 1, the lobe had overfilled the 50-m-wide notch. For additional information on the Mount St. Helens extrusion, see the excerpts from the monthly Scientific Event Alert Network (SEAN) in the news section of this issue. (Photograph taken by Tom Casadevall, U.S. Geological Survey, Cascade Volcano Observatory, 5400 MacArthur Blvd., Vancouver, WA 98661; photograph courtesy of SEAN.)

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Lokesh Chaturvedi has been working full time with the Environmental Evaluation Group (EEG) of New Mexico since June 1982 to evaluate the site characterization work being conducted by DOE for the WIPP project. His involvement with WIPP began in 1979 when he started working as a consultant to EEG while he was a professor of geological engineering at New Mexico State University. He has a Ph.D. degree in geological sciences from Cornell University, M.S. in civil engineering from Purdue University, M.Sc. in geology from Roorkee University (India) and a bachelor's degree in geology, physics, and math from the University of Rajasthan (India). He has taught at Michigan Tech., Indian Institute of Technology, City University of New York, and New Mexico State University. Before becoming involved in the site characterization for nuclear waste disposal, his research work was in geothermal hydrology and remote sensing for which he has conducted field work in Iceland, Himalayas, and the western United States.

Kenneth R. Rehfeldt worked as a geohydrologist with the Environmental Evaluation Group during 1982–1983. He reviewed the hydrologic investigations for WIPP and independently analyzed a large amount of data which resulted in EEG's recommendations to DOE for additional hydrologic work. He received his M.S. in hydrology from the New Mexico Institute for Mining and Technology and his bachelor's degree in geology from the University of Wisconsin-Madison. He is currently working toward his Ph.D. at the Massachusetts Institute of Technology and is conducting research in defining the controlling mechanisms of solute transport processes in groundwater.

increased precipitation in the north and northwestern parts of North America. Evaporation also increased by 15–20% in the northern and western parts of the continent and by 11% worldwide, similar to the percentage increases for rainfall. Runoff increased over the northwestern and extreme southwestern parts of the continent by as much as 20–60%.

In order to minimize computer time, the researchers divided North America's weather into large "grid areas" approximately 27,000 km square, and so were not able to consider the specific effects of different topographical features on the weather. As a result, they point out, the current computer model is able to accurately recreate existing climatic conditions "only at a very aggregate level."

But the modeling work needs to continue, say the authors of the report, which is entitled "Potential Climatic Impacts of Increasing Atmospheric CO<sub>2</sub> with Emphasis on Water Availability and Hydrology in the United States." "This report, in effect, presents a methodology for estimating the hydrologic impact of increased atmospheric CO<sub>2</sub>," they state in their introduction, "and should be looked upon as a first approach to a complex problem."—TR

## Science Policy Studied in Congress

The House Science and Technology Committee will initiate a comprehensive study of science policy in the United States. Although the study will not formally begin until January, when the 99th Congress convenes for its 2-year term, a newly appointed task force has begun to develop the agenda for the committee's work and has begun to prepare background information for the study.

Dou Fugua (D-Fla.), chairman of the Science and Technology Committee, said that the health and vitality of American science unquestionably has been a major factor in the strong performance of the American economy over the last 35 years. However, the committee is concerned that present policies and practices may not be fully adequate to the new environment facing U.S. science in the coming decades.

Among the issues that are expected to be on the agenda are the institutional framework for the support and conduct of scientific research; the training and education of young scientists; methods of funding research; and the overall funding levels for science.

The task force, to be composed of approximately 15 members, will be chaired by Fugua. News (cont. on p. 460)





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## Geophysical Events

This is a summary of *SEAN Bulletin*, 9(6), June 30, 1984, a publication of the Smithsonian Institution's Scientific Event Alert Network. The complete bulletin is available in the microfiche edition of *Eos* as a microfiche supplement or as a paper reprint. For the microfiche, order document E84-007 at \$2.50 (U.S.) from AGU Fulfillment, 2000 Florida Avenue, N.W., Washington, DC 20009. For the paper reprint, order *SEAN Bulletin* (giving volume and issue numbers and issue date) through AGU Separates at the above address; the price is \$3.50 for one copy of each issue number for those who do not have a deposit account, \$2 for those who do; additional copies of each issue number are \$1. Subscriptions to *SEAN Bulletin* are available from AGU Fulfillment at the above address; the price is \$18 for 12 monthly issues mailed to a U.S. address, \$28 if mailed elsewhere, and must be prepaid.

## Volcanic Events

Merapi (Java): Explosions, nuees ardentes, lahars: 1000 evacuated.

Tinakula (Santa Cruz Is.): Tephra ejection; W flank submarine cone recognized.

Rabaul (New Britain): Seismicity declines; deformation continues.

Mamam (Bismarck Sea): Strombolian activity; debris avalanches.

Langila (New Britain): Occasional ash emission; seismicity weak.

Campi Flegrei (Italy): Seismicity and uplift continue; socio-economic impact discussed; increased submarine fumarolic activity.

Etna (Italy): Explosions and lava production continue from SE Crater; central crater explosions.

Kilauea (Hawaii): Phases 20-22, highest lava fountains of 1983-84 eruption.

Mount St. Helens (Washington): New lobe extruded into notch in dome's W flank (see cover photograph).

Atmospheric effects: Atmospheric turbidity over Japan declines gradually from late 1982 to early 1983 peak; lidar shows persistent aerosols.

Merapi Volcano, Java, Indonesia (7.5°S, 110.4°E). All times are local (= UT + 7 hours).

The quoted material below is excerpted from a report by Adjat Sudrajat.

Merapi erupted June 15 between 0215 and 0600, accompanied by nuees ardentes that extended 7 km down rivers (the Batang, Sebeg, and Krasak) on the SW side of the volcano. An eruption plume rose to 6 km height and caused ashfall in Blantian, Arabarua, and Semarang. Approximately 60 km north of the volcano. The eruption was accompanied by detonations. The first explosion was followed by a milder eruption producing a plume to 2 km height and a nuee ardente to 6 km distance at noon. The frequency of nuees ardentes progressively decreased until the morning of June 16. No eruptions were observed the following day.

Lahar material estimated to exceed 4 x 10<sup>6</sup> m<sup>3</sup> along the Bedeng, Krasak, and Putih rivers may threaten Magelang city (population about 125,000).

Newspapers reported ashfall at Magelang (30 km NW of Merapi) and Salatiga (35 km NE of the volcano). Visibility near Salatiga was limited to 10 m and more than a centimeter of ash covered roads, slowing traffic. More than 2 cm of ash fell at Solo (15 km E of the volcano), and ashfall was reported at Cilacap, on the coast 160 km SW of Merapi.

"Seismographs detected a progressive increase in seismicity from four counts per day June 8 to 59 on June 12. A warning was issued June 13 and the evacuation of 1,000 persons from forbidden zone section VI (Kendren village) was immediately implemented. The long-term warning system was tested again to be sure that it was operational. The eruption was preceded by an intense lava avalanche on June 13 that caused a nuee ardente d'avalanche (nuee ardente of Merapi type)."

Information Contact: Adjat Sudrajat, Director, Volcanological Survey of Indonesia, Diponegoro 57, Bandung, Indonesia; The Jakarta Times, Jakarta, Indonesia.

Tinakula Volcano, Santa Cruz Islands, SW Pacific Ocean (10.47°N, 167.75°E).

The following is from the cruise report of the USGS research vessel S. P. Lee, engaged in multichannel seismic profiling in the Vanuatu and Solomon Islands areas.

"On June 3, Tinakula could be seen 'smoking' in the distance, some 25 km away. As the ship approached the island, large billowing clouds of steam were observed. At approximately 2-hour intervals, a large billowing dark gray ash-laden plume was observed rising to several kilometers above the island. The S. P. Lee passed within 400 m of Tinakula along the N side of the island. Rumbling sounds could be heard from within the active vent, immediately north of the central crater. Boulder-size rocks were ejected from the vent and were still steaming as they rolled and skipped down the steep scree slope into the sea. At least a dozen or so of the boulders, and much more material of cobble size, were seen being thrown from the vent every minute. Much of this debris was accumulating on the scree slope.

"Geophysical data collected by the S. P. Lee showed that another volcanic cone is present, about 90 m beneath the surface of the water, some 5 km W of Tinakula. It has a sharp steep flank, and appears from its morphology to

be active. This volcano or volcanic vent has not been identified before and is not on any bathymetric map."

Tinakula's last reported eruption occurred September 6 to December 11, 1971. Interim explosive activity built a small summit cone; incandescence blocks rolled down the volcano's flanks; and a slow-moving lava flow extended about 300 m down the NW flank.

Information Contacts: A. Macfarlane, Director, Department of Geology, Mines, and Rural Water Supplies, GPO, Port Vila, Vanuatu; H. G. Greene, Branch of Pacific Marine Geology, USGS, 345 Middlefield Road, Menlo Park, CA 94025.

Mt. St. Helens Volcano, Cascade Range, S Washington, USA (46.20°N, 122.18°W). All times are local (= UT - 7 hours).

The quoted material is from Peter Otway.

A new lobe in the composite lava dome began to emerge in mid-June. Its location, on the W flank, was within the notch that was the source of explosions on May 14, 26, and 27 and June 6 (see *SEAN Bulletin*, vol. 9, nos. 4-5) as well as a similar event on June 7 (see below). Growth of the vent flank lobe had stopped by July 1. Accelerating deformation on the dome's north side was measured in late June and early July, but rates of deformation began to decrease after several days and no lava reached the surface.

A small explosion occurred June 7 at 1720 when airline pilots observed an ash cloud that rose to about 9 km altitude. Increased water

flow from the crater began at about 1740, and a mudflow about 3 m wide and 15 cm deep reached Spirit Lake (5.5 km from the crater) at 1802.

Beginning June 14-15, the number of earthquakes at the nearest crater seismometer (Yellow Rock) increased from about 20-30/day to 50-60/day, and the number of surface (rockfall) events began to increase June 16-17. The small new lobe was first seen during the afternoon of June 17 (by gas monitoring aircraft pilot Al Maris), and was more clearly visible during an overflight at 2230 that night.

When first tracked, on June 18, the leading edge of the lobe, was moving down slope at over 30 m/day, although rates near the extrusion site were at least 50% higher. By June 22, movement of the leading edge had slowed to 13 m/day, and by June 25, to 6 m/day. When next observed, on July 1, the flow had stopped. The lobe is 60 m wide at its maximum and 150 m long. Its volume is estimated to be on the order of 0.2 x 10<sup>6</sup> m<sup>3</sup>.

Targets on the north side of the dome moved north/northeast at rates which steadily

increased to 60 mm/day by June 25. Movement then swung to the north as rates accelerated rapidly to peak at 0.8 m/day between June 30 and July 2. By July 10, the rates had fallen to 20 mm/day. This suggests that an intrusive event occurred late in June at a site at least 100 m east of the mid-June extrusion.

Rates of SO<sub>2</sub> emission ranged from 15 to 35 metric tons per day (near the detection limit) during the first 2 weeks of June. The net measurement, during the morning of June 18, was 105 metric tons per day. Rates averaged about 100 tons per day through July 1, but had dropped to 40 tons per day by July 6.

Information Contacts: Steven Brantley, Tom Casadevall, Eliot Endo, Clint Mullins, Chris Newhall, and Peter Otway, USGS Cascades Volcano Observatory, 3400 MacArthur Blvd., Vancouver, WA 98661; Robert Norris, Geophysics Program, University of Washington, Seattle, WA 98195.

## Meteoritic Events

Fireballs: Arizona, Arkansas, Iowa, USA.

## Earthquakes

Date	Time, UT	Magnitude	Latitude	Longitude	Depth of Focus	Region
June 24	1117	6.5 M <sub>L</sub>	18.03°N	69.33°W	shallow	S of Hispaniola

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Applicants should send resume and names of three references to Robert A. Philpney, Chairman, Department of Geological and Geophysical Sciences, Princeton University, Princeton, N.J. 08544.

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## Correction

The following corrections to the article, "The Lunar Core and the Origin of the Moon," by Horton E. Newsom, which was published in *Eos*, p. 369, May 29, 1984, were not included in the published version. A corrected version of the article will appear in the shelf edition of *Eos*.

1. The author's name was misspelled. It should be Horton E. Newsom.

2. The author's current address is Institute of Meteoritics and Department of Geology, University of New Mexico, Albuquerque, NM 87131.

3. In Table 3, P/Nd C1 value is 2,100.

4. Among some references not included with the published version was a reference for A. E. Ringwood and S. E. Kesson, Basaltic magnetism and the bulk composition of the moon, 2, Siderophile and volatile elements in moon, earth and chondrites: Implications for lunar origin, *Moon*, 16, 425-464, 1977.

5. Dr. Newsom would like to acknowledge A. E. Ringwood, whose additional suggestions were not included in the published version. The modern concept of a terrestrial origin for the moon remains indelibly stamped with Ringwood's name.

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